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Soil fertility in slash and burn agricultural systems in central Mozambique

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27 **Abstract**

28

29 Slash and burn is a land use practice widespread all over the world, and nowadays it is formally
30 recognized as the principal livelihood system in rural areas of South America, Asia, and Africa. The
31 practice consists of a land rotation where users cut native or secondary forest to establish a new crop
32 field and, in some cases, build charcoal kilns with the cut wood to produce charcoal. Due to several
33 socio-economic changes in developing countries, some scientists and international organizations have
34 questioned the sustainability of slash and burn since in some cases, crop yield does not justify the soil
35 degradation caused. To estimate the soil quality in agricultural and forest soils at different ages of the
36 forest-fallow period (25, 35, and 50 years), this survey investigated rural areas in three locations in
37 Manica province, central Mozambique: Vanduzi, Sussundenga, and Macate. Soil profiles were
38 trenched and sampled with a pedological approach under crop fields and forest-fallow. The
39 chronosequence was selected to test the hypothesis that the increase in forest-fallow age causes an
40 improvement of soil fertility. Results highlighted discrete variations among locations in mineralogy,
41 Al- and Fe-oxyhydroxides, sand, silt, pH, total organic carbon, humic carbon, total nitrogen, available
42 phosphorous, chloride, nitrate, fluoride, and ammonium. Few differences in mineralogy, Fe-
43 oxyhydroxides, available P, chloride, and nitrate were detected between crop fields and forest-fallow
44 within the same location. Such differences were mostly ascribed to intrinsic fertility inherited from
45 the parent material rather than a longer forest-fallow period. However, physicochemical soil property
46 improvement did not occur under a forest age of 50 years (the longest forest-fallow considered),
47 indicating that harmonization of intrinsic fertility and agronomic practices may increase soil organic
48 matter and nutrient contents more than a long forest-fallow period.

49

50 **Keywords:**

51 Soil fertility, sustainability, slash and burn agriculture, tropical soils, forest-fallow

52 1 Introduction

53 *Slash and burn* agriculture is widespread in tropical and sub-tropical regions all over the world
54 (Li et al., 2014; Mukul and Herbohn, 2016) and occupies an important role in subsistence and cultural
55 identity (e.g., Brady, 1996; van Vliet et al., 2012; Mukul and Herbohn, 2016). Paleoecologists have
56 observed that, in the Amazonian basin, the first adoption of slash and burn practice dates to the pre-
57 Columbian population (Arroyo-Kalin, 2012). On a global scale, an estimation by Hauser and
58 Norgrove (2013) reported that 36 million km² of land was under slash and burn system, about 1/3 of
59 the global soil resource. In terms of the population involved, FAO/UNDP/UNEP (2008) roughly
60 estimated that over 500 million people were using slash and burn in 1982; however, these data are
61 difficult to recollect (Mertz et al., 2009) and must have certainly changed in the last 40 years. Slash
62 and burn consisted, and still consists, in cutting, burning, and farming different forest areas in rotation,
63 often with the cut stems and branches used to create charcoal kilns (3-6 per hectare), while stubbles
64 are burnt in situ (e.g., Brady, 1996; Brown, 2006; Riahtam et al., 2018). In some areas, the charcoal
65 produced by charcoal kilns is sold in local markets (Kabisa and Ncheengamwa, 2020), while the very
66 small fragments are incorporated with ashes into the soil by a basic plough to obtain a small increment
67 of soil fertility and crop production (e.g., Rumpel et al., 2006; Gay-des-Combes, 2017; Selvalakshmi
68 et al., 2018). The cleaned-up area is proportioned to family food requirements, and users continue to
69 exploit it until the productivity goes below the sustenance threshold (Jakovac et al., 2016); after that,
70 they move to another forest spot. The once cultivated areas are left to long fallow, with progressive
71 regeneration of spontaneous vegetation and development of a secondary forest. The latter will be cut
72 down again to begin a new cropping cycle (Gonçalves Lintemani et al., 2019; Hauser and Norgrove,
73 2013). This system has gone on for centuries in most of the suitable regions where peasants are pushed
74 by environmental conditions to apply it for obtaining essential goods (e.g., Tschakert et al., 2007;
75 Edivaldo and Rosell, 2020). In the last 4-5 decades, exploitation intensity has changed in slash and
76 burn systems concurrently with *i*) the demographic increment and the subsequent need for more
77 cultivable land (Kilawe et al., 2018), *ii*) economic policies pushing to convert traditional cultivations

78 into cash crops (Vongvisouk et al., 2014; Wood et al., 2016; Kilawe et al., 2018), and *iii*) the necessity
79 to obtain products to sell in local markets to achieve a small income (Ickowitz, 2011). Therefore,
80 while once the area was left to fallow to recover to forest (hereinafter referred to as forest-fallow)
81 over periods of 50-100 years, in the last decades the forest-fallow period has been progressively
82 shortened to 10-20 years (Juo and Manu, 1996). The forest-fallow length is considered one of the
83 main limitations of this land use system since it is expected that the soil needs many decades to restore
84 fertility after cultivation. Fachin et al. (2021) studied revegetated soils in Paraná (Brazil) submitted
85 to slash and burn system following a fallow chronosequence from one month to 12-years and found
86 that chemical properties did not directly increase with the fallow age and development. Other issues
87 raised by the shortened forest-fallow period, which is a form of deforestation (Dirac Ramohavelo,
88 2009; Mukul and Herbohn, 2016), are the reduced capability of carbon sequestration and storage
89 (Kotto-Same et al., 1997), the erosion and loss of soil nutrients (Runyan et al., 2012; Thomaz, 2013),
90 and the reduction of forest products like charcoal, medicines, fruits, nuts, and artisanal materials
91 (Junsongduang et al., 2013). All over the world, many authors studied these effects from various
92 points of view by investigating the soil physicochemical properties (e.g., Alegre and Cassel, 1996;
93 Fachin et al., 2021; Rumpel et al., 2006), plant biodiversity, and vegetation dynamics (e.g., De Wilde
94 et al., 2012; Randriamalala et al., 2019), the release of greenhouse gases (Davidson et al., 2008;
95 Dhandapani and Evers, 2020), the chemical composition of charcoal, the stable soil C stocks
96 (Selvalakshmi et al., 2018), and the microbiota composition (Aboim et al., 2008; Kukla et al., 2019).
97 For most of these investigations, slash and burn is considered not sustainable any longer from social
98 and environmental points of view. However, in addition to the shortened forest-fallow, other factors
99 should be considered as limiting the sustainability of this practice: *i*) extreme climatic conditions; *ii*)
100 quartz-rich soil parent materials and generally edaphic conditions (Kleinman et al., 1995); *iii*)
101 agroecological settings i.e., the number of fallow/cropping cycles, fallow species succession, use of
102 fertilizers, soil management (e.g. Styger et al., 2007; Mertz et al., 2008; Coomes et al., 2017); *iv*)
103 technical efficiency at farming, infrastructure, and policy level (Binam et al., 2004).

104 Because of the crucial role of slash and burn in many economies of developing countries, soils of
105 three locations in central Mozambique where the slash and burn system has been historically practiced
106 were investigated. The locations were selected based on the different length of the forest-fallow and
107 cultivation period, obtaining a chronosequence of 25, 35, and 50 years for the forest-fallows and a
108 chronosequence of 1, 2, and 16 years for the crop fields. In each location, both crop field (CF) and
109 forest-fallow (FF) were considered.

110 This work aimed to evaluate the impact of the slash and burn system on soil properties and fertility.
111 Specifically, the morphological, mineralogical, and physicochemical (including soil texture,
112 extractable Al- and Fe-oxyhydroxides, and water-soluble ions) soil properties were investigated to
113 assess if and how several years of forest-fallow can influence the fertility of the revegetating areas.
114 Two hypotheses were tested: 1) the longer the forest-fallow, the higher the recovery of the soil fertility
115 as a result of organic matter accumulation; 2) the soils under FFs are more fertile than those under
116 CFs.

117

118 **2 Materials and methods**

119 *2.1 Characteristics of the locations and selection of the study areas*

120 Details about agroecological, geological, and social background are reported at Point 1 of
121 Supplementary Materials. Following the Köppen-Geiger climate classification, the climate of the R4
122 agroecological region is humid sub-tropical, with a warm temperate climate with a dry and cold
123 season from April to September and a hot and rainfall season from October to March (Kottek et al.,
124 2006; Belda et al., 2014), with frequent storms and cyclones. The soil parent rock at Vanduzi and
125 Sussundenga was a granitoid rock (possibly gneissic-granite) belonging to the Barue Magmatic Arc
126 (Wijnhoud, 1997; Chaúque et al., 2019), while at Macate it was a migmatitic paragneiss belonging to
127 the supracrustal rocks of the Chimoio group (Chaúque et al., 2019) (see Supplementary materials).

128 With repeated surveys in the three districts, it was ascertained that in some fields the remainders of
129 previous charcoal kilns were even more than 20 per hectare and, in some cases, charcoal kilns had

130 been superimposed one over the other several times. As witnessed by the presence of charcoal kiln
131 residues, all the surveyed forests were areas left to long fallow after cropping in which a *miombo*
132 forest developed, as also reported by Montfort et al. (2021). Details on *miombo* are presented at Point
133 2 of Supplementary materials. In all cases, it was also ascertained that conditions of the forest-fallow
134 were rather poor in terms of the number and species of trees (Table 1); thus, in interviews, farmers
135 claimed that both biodiversity and tree density were higher 60-70 years ago. In each district, a study
136 area was selected following information obtained mainly by interviews for Vanduzi and Macate,
137 while for Sussundenga information was also retrieved from documents provided by the Research
138 Station at IIAM (Mozambican Institute of Agricultural Research). The selected study area included
139 both CF and FF representative of the R4 agroecological region to form a FF chronosequence: 25 years
140 at Vanduzi, 35 years at Sussundenga, and ≈ 50 years at Macate. The CFs close to the FF areas were
141 cultivated with annual and/or pluriannual crops established with the slash and burn system.

142 The main characteristics of the study areas are reported in Table 1 and briefly below.

- 143 • *Vanduzi* study area was on a gentle slope (3%), at an altitude of 658 m. CF was established one
144 year before the survey by shallow ploughing (about 10 cm) to be devoted to vegetable
145 garden/banana orchard managed with manual hoeing and irrigation, without fertilizer
146 application. FF was a < 25 years old *miombo* moderately vigorous with crown-gaps and renewal,
147 whose regrowth was disturbed by occasional grazing and low-intensity fires, and used to obtain
148 firewood, timber, and charcoal. The main tree species was *Mangifera indica* L. (about 20% of
149 tree abundance), remainders of an abandoned crop field with sparse mango plants, together with
150 species typical of *miombo* (about 80% of tree abundance).
- 151 • *Sussundenga* study area was in an area with a gentle slope (2-4%), at an altitude of 649 m. CF was
152 established two years before the survey after shallow ploughing with cow traction (8-10 cm).
153 The field was cultivated with maize, without irrigation and fertilizer application, and sowed after
154 the burning of plant residues and shallow ploughing with cow traction. The average maize grain
155 yield was estimated to be 1.0-1.1 Mg ha⁻¹ y⁻¹. Documents attested that FF was 35 years old

156 *miombo*; it appeared moderately vigorous with the presence of gaps and renewal and was
157 occasionally used as a source of timber.

158 • *Macate* study area was in an area with a gentle slope (3%), at an altitude of 555 m. CF was
159 established after shallow ploughing with cow traction (8-10 cm) and continuously cultivated with
160 maize for 16 years with manual hoeing and without irrigation and fertilizer application. Yield
161 roughly ranged between 1.1 to 1.5 Mg ha⁻¹ y⁻¹. Crop residues were commonly used for mulching
162 and, only at the end of the 2017 wet season, a controlled crawling fire was applied to reduce the
163 invasive *Mucuna pruriens* DC. FF was ≈50 years old *miombo*, vigorous, without gaps but with
164 renewal. It has been and is used as a source of timber, firewood, and game.

165 To prove the age of the forest-fallow, with the counting of tree rings being useless, we ascertained
166 that the average tree diameters of the ubiquitous *Brachystegia spiciformis* trees of Macate (33 cm)
167 was higher compared with that of Sussundenga (26 cm) and Vanduzi (16 cm) trees.

168

169 2.2 Soil sampling procedure

170 In March 2017, in each of the three study areas, a geomorphological and soil survey was run to
171 select sampling sites with similar exposure and slope (Table 1). In doing this, several mini-pits and
172 auger holes were opened before choosing the best position where to dig the soil profiles. Soil profiles
173 were trenched (0.8 m of width) in both CFs and FFs, all within a surface of ≈1 ha (Fig. 1).

174 In the FFs, profiles were opened at 1-1.5 m downslope from the trunk of one of the biggest
175 *Brachystegia spiciformis* Benth. trees, where the stem influence was considered null. In the CFs,
176 profiles were opened in the middle of their extension. The maximum distance between CF and FF
177 was about 60 m at Macate and Sussundenga, while at Vanduzi FF and CF sites were at a distance of
178 about 700 m. In all cases, sampling sites were at least 30 m from the rather sharp CF-FF transition, a
179 distance that was considered sufficient to avoid considerable edge effects. As a replicate, a second
180 soil survey was made in November 2017, in which soil profiles located a few meters from the previous
181 ones were opened. As a whole, 12 profiles were sampled (3 locations per 2 land uses per 2 replicates).

182 The soil profiles were morphologically described per Schoeneberger et al. (2012) and about 4 kg of
183 soil sample from each genetic horizon were collected. During the field activities, the collected
184 samples were maintained inside a portable fridge and, once in the laboratory, they were air-dried,
185 sieved at 2 mm to remove the skeletal particles, and maintained at 4°C for a maximum of one week
186 before the analyses.

187

188 2.4 *Soil analyses*

189 Soil mineralogical assemblage was evaluated by X-ray diffractometry on manually compressed
190 powdered samples by using a Philips PW 1830, which produced a Fe-filtered Co K α 1 radiation,
191 operating at 35 kV and 25 mA. The identification of the minerals was done based on their
192 characteristic peaks (Brindley and Brown, 1980; Dixon and Schulze, 2002). A semi-quantitative
193 assessment of the mineralogical assemblage was obtained by estimating the area of the primary peaks
194 by multiplying the peak height by its width at half-height (Cocco et al., 2015).

195 The particle-size distribution was determined by the pipette method (Day, 1965) after the dissolution
196 of organic cements by NaClO at pH 9 (Lavkulich and Wiens, 1970). The sand (2-0.05 mm) was
197 separated by wet sieving, while silt was separated from clay by column sedimentation at 19-20°C air
198 temperature. The pH was determined potentiometrically in water (1:2.5 solid:liquid ratio), using a
199 combined glass-calomel electrode immersed into the suspension. The content of total organic C
200 (TOC) was estimated by K-dichromate digestion, heating the suspension at 180°C for 30 min
201 (Allison, 1965), and the humic C (HC) was determined by the Walkley-Black method (Nelson and
202 Sommers, 1996). The total nitrogen (TN) was determined by the Kjeldahl semi-micro method and
203 the potentially plant-available phosphorous (AvP) was estimated according to Olsen et al. (1954).

204 The amounts of Al and Fe forming pedogenic oxyhydroxides were estimated by extraction with
205 dithionite-citrate-bicarbonate (DCB) treatment (Jackson, 1958), using 30 ml of mixed solution plus
206 1 g of Na-dithionite to treat 2 g of soil sample; the extraction was repeated two times and then washed

207 twice. The total extract was filtered with an ashless Whatman 42 filter paper, and the Al and Fe in
208 solution were determined by Optical Emission Spectrometer – Optima 8300.
209 Soluble ions were extracted by distilled water (1:10 solid:liquid ratio). After gentle stirring for 1 min,
210 the suspension was left to rest for few minutes and the supernatant was filtered with ashless Whatman
211 42 filter paper. The concentration of anions (chloride, fluoride, sulphate, nitrate, phosphate, nitrite,
212 bromide, acetate, oxalate, and formate) and cations (calcium, magnesium, potassium, sodium,
213 ammonium, and lithium) was determined in the solution by a Dionex ICS-900 Ion System
214 Chromatograph equipped with IonPac AS23 column for anions and IonPac CS12 column for cations,
215 using a 0.5 M Na-bicarbonate solution as eluent. The concentration of bicarbonate ions was obtained
216 from the difference between the summation of cations and the summation of anions.

217

218 2.5 *Statistical analysis*

219 For each horizon, a single determination was performed for mineralogy, particle-size
220 distribution, pH, TOC, HC, TN, AvP, and total extractable (pedogenic) Fe and Al. For the soluble
221 cations and anions, two extractions per sample were obtained each, and the two values averaged to
222 obtain more reliable results. R program (1.3.1093, R Core Team, 2014) was used for statistical
223 analysis, that was run for soil physicochemical properties among soil profiles. The ANOVA at 5%
224 significance was used to assess those physicochemical results obtained by the two surveys (Tables
225 S1, S2, S3, and S4 of Supplementary materials). Therefore, the results of the two surveys were
226 considered replicates, and the results of each horizon were used to calculate the weighted average
227 based on the thickness of each horizon; the standard deviation was also calculated. Principal
228 Component Analysis (PCA) [FactoMineR and factoextra R packages] for the entire dataset was
229 impossible to compute since the cumulative variance percentages were > 60% over the third
230 dimension. Therefore, PCA was performed for particle-size distribution, pH, TOC, HC, TN, AvP,
231 and extractable Al and Fe in soil profiles (Fig. 2) to assess the structure of the main variables in the
232 three study areas and in CF and FF. To further analyse the results, ANOVA was applied to enhance

233 significant differences in soil profile properties among CFs and FFs from the three locations, and
234 between CF and FF within each location. Data were tested for normality and homoscedasticity by
235 performing the Shapiro-Wilk statistical test (stats R package) and by the Levene's test (car R
236 package), both at 5% of significance level, respectively (Tables S5, S6, S7, S8, and S9 of
237 Supplementary materials). Box and Cox transformation was used to transform the data in case they
238 were not parametric (Meloun et al., 2005). If the transformed data were normally distributed, a post-
239 hoc Tukey's Honest Significant Difference (HSD) test with $P \leq 0.05$ was used to compare the means.
240 When normality was not respected, the Kruskal-Wallis test was applied. In the case of
241 heteroscedasticity, the Welch one-way ANOVA test was performed. ANOVA tests were deemed
242 significant when $P \leq 0.05$.

243

244 **3 Results**

245 *3.1 Soil morphology and mineralogy*

246 In all the locations (Vanduzi, Sussundenga, and Macate), the soils belonged to the order of
247 Oxisols according to Soil Survey Staff (2014) or Ferralsols according to IUSS (2015), due to the
248 presence of A (umbric) and Bo (oxic) horizons (Table 1). In general, A and Bo horizons showed red-
249 yellowish colour and a good degree of aggregation made of blocks generally coarser in the A than in
250 the Bo horizons. Such good state of aggregation, the coarse texture (mainly sandy loam), and the
251 absence of redoximorphic features indicated these soils as well-drained and, consequently, with low
252 to moderate water-holding capacity (Agrawal, 1991; Suzuki et al., 2007). Roots were rather abundant,
253 from very fine to coarse in size, in the A and Bo horizons under FFs, while under CFs root density
254 and size decreased along the profile and were absent in the cultivated soil of Macate. The A horizons
255 under CFs always showed a charcoal content of $< 1\%$, while under FF charcoal fragments were found
256 rarely in the subsurface horizons. Two of the three FFs presented organic horizons, identified as Oi
257 horizons at Sussundenga and as Oi and Oe&Oa horizons at Macate. Only under CF at Vanduzi and
258 FF at Sussundenga, soil tunnelling due to termites was found.

259 In all the soils mineralogical composition was dominated by quartz, with the highest contents at
260 Sussundenga (90-95%) and the lowest at Macate (67-79%). A major variability in mineralogical
261 composition was observed in the Macate soils, where plagioclases and kaolinite were present in
262 higher quantities than in the Vanduzi and Sussundenga soils. Vanduzi showed the highest content of
263 2:1 clay minerals under CF and the lowest content of kaolinite under both land uses. Within each
264 location, the only difference between land uses was observed at Macate, where a slightly higher
265 content of 2:1 clay minerals was found in the FF soil than in the CF soil (Table 2).

266

267 3.2 *Physicochemical properties, including extractable (pedogenic) Al and Fe*

268 The PCA clearly separated soils of Macate from those of the other study areas mainly for the
269 drivers of sand, silt, TOC, HC, TN, AvP, and extractable (pedogenic) Al and Fe variables, which
270 were explained by the first component (PC1) for 47.5% and by the second component (PC2) for
271 24.3% of the variance (Fig. 2). In all the soils the particle-size distribution revealed a predominance
272 of sand, followed by clay and silt. Both CF and FF soils of Sussundenga and FF soil of Vanduzi
273 displayed higher content of sand than all the soils of Macate (Fig. 3A). The FF soil of Macate showed
274 a major abundance of silt compared with FF soils of Vanduzi and Sussundenga (Fig. 3B). No
275 significant difference was observed between CF and FF soils within each location (Fig. 3A, B, C).

276 The Vanduzi soils were slightly acid, with values ranging from ≈ 6.3 to 6.8 (Fig. 3D). As expected,
277 the contents of TOC, HC, and TN were low on absolute value but resulted higher in the CF soil of
278 Macate than in those of Vanduzi and Sussundenga (Fig. 3E, F). Thus, while among FF soils no
279 significant difference was observed for TOC and HC, TN displayed the highest concentrations in both
280 CF and FF soils of Macate (Fig. 3G). Since the soil deficiency threshold for the adopted method is
281 considered to be 23 mg kg⁻¹ (Cardelli et al., 2017), the AvP content was always very low, with the
282 highest content in the FF soil of Vanduzi, where it was also higher than in the CF soil (Fig. 3H).

283 The total extractable Al ranged from 1.1 to 7.1 g kg⁻¹ in all the soils, while Fe varied from 3.3 to 40.4
284 g kg⁻¹ (Table 3). The soils of Macate showed the highest concentrations of extractable Al and Fe in

285 both CF and FF soils. In addition, only at Macate, the FF soil showed a higher concentration of
286 extractable Fe than the CF soil (Table 3).

287

288 3.3 *Water-soluble anions and cations*

289 Formate and lithium were always below the detection limits, while bromide was detected in a
290 very small amount ($\approx 1 \mu\text{eq kg}^{-1}$) only in the Ap horizon of Macate CF soil (Tables S10, S11).
291 Significant ($P < 0.05$) differences for anions were detected only for chloride, fluoride, and nitrate,
292 whereas among cations only for ammonium (Fig. 4). Chloride concentration was similar in all the
293 soils under CF (mean $496 \mu\text{eq kg}^{-1}$), while under Macate FF soil it showed a higher value than in the
294 Sussundenga soils (400 and $254 \mu\text{eq kg}^{-1}$ respectively). Chloride also differed between land uses in
295 the Vanduzi soils, with higher concentrations in CF than in FF soils (408 and $318 \mu\text{eq kg}^{-1}$
296 respectively) (Fig. 4A). The nitrate concentration was higher at Macate than at Sussundenga for CF
297 soils (342 and $58 \mu\text{eq kg}^{-1}$ respectively), and it was higher at Macate than at Vanduzi for FF soils
298 (804 and $5 \mu\text{eq kg}^{-1}$ respectively) (Fig. 4B). Fluoride was the highest in the soils of Vanduzi and
299 Macate, regardless of the land use (8 and $9 \mu\text{eq kg}^{-1}$ respectively) (Fig. 4C). Ammonium was higher
300 in the soils of Macate and Sussundenga than in those of Vanduzi, independently of the land use (200 ,
301 215 , and $20 \mu\text{eq kg}^{-1}$ respectively) (Fig. 4D).

302

303 4 Discussion

304 4.1 *Impact of the forest-fallow length on soil properties and fertility*

305 4.1.1 *Soil morphological and physical properties*

306 The predominant reddish colour of the soils and the abundance of quartz and sand particles
307 underlined the strong weathering processes faced by the studied soils, where the more easily
308 weatherable minerals in the upper soil meter have been removed by alteration. Similar results were
309 reported by Sá et al. (1972) and FAO (1982) and attributed to the lateritization process that is
310 responsible for the development of Oxisols (Van Wambeke et al., 1983; Soil Survey Staff, 2015).

311 This process is also responsible for the formation of a good soil structure because of the progressive
312 increment of Al- and Fe-oxyhydroxides that act as cementing agents (Igwe et al., 2009; Krause et al.,
313 2020). The small mineralogical variability among locations betrayed the different parent materials
314 from which the soils have developed. In fact, the lowest quartz and the highest kaolinite contents in
315 the soils of Macate were probably due to the composition of their parent rock, a migmatitic paragneiss,
316 which is a lithology generally richer in fine-grained clastic sediments than the granitoid rocks of the
317 other two sites. Crystalline rocks like granites and gneisses may contain intercalations of sedimentary
318 silty or clayey beds (Gray et al., 2016), and this could be the explanation for the different particle-
319 size distributions of these soils. The occurrence of tunnelling in the CF soil of Vanduzi and in the FF
320 soil of Sussundenga was ascribed to attempts of termites to colonize these soils; as a matter of fact,
321 no termite nests were observed in the hectare of the surface under study, nor in the close surroundings.
322 Therefore, the small morphological and physical differences among locations appeared related to the
323 composition of the parent rocks rather than to the uncultivated period duration, expressed by the
324 forest-fallow age, or the cultivation period.

325

326 4.1.2 Soil chemical properties

327 The relatively high pH values in the Vanduzi soils were ascribed to the higher presence of 2:1 clay
328 minerals (mainly vermiculite), which offer a much higher buffering capacity than the other minerals
329 present (Abate and Masini, 2005; Malandrino et al., 2006; Abollino et al., 2008). Because of this, the
330 distinct parent material and related mineralogical assemblage appeared to be the main cause of the
331 different soil pH values. TOC and HC were present in very low amounts, in line with the contents
332 reported by Rafael et al. (2018) for Mozambican soils. However, since both TOC and HC contents
333 were similar in the FF soils and were the highest in the Macate CF soil, deductions were that:

334 *i)* in this environment, a 50-years forest-fallow is not able to enrich the FF soils with organic matter.

335 Studying the soil characteristics after slash and burn, Montfort et al. (2021) found that organic
336 carbon stock in the upper 30 cm soil was similar in 20 and 25 years old *miombo* forests and mature

337 woodland due to the rapid vegetation regeneration and that disturbances typical of slash and burn
338 (biomass removal, fire, and soil tillage) can decline this content. Moreover, a study conducted by
339 Williams et al. (2008) in arid Mozambican *miombo* forests reported no identifiable changes in soil
340 organic C accumulation along a chronosequence in the re-growing miombo (for a maximum age
341 of 30 years) due to the extremely slow input of organic matter to the soil; and that,
342 *ii*) for the higher TOC and HC contents in CF soils of Macate, it occurred for the soil cultivated for
343 the longest time (16 years) compared to the soils of Vanduzi and Sussundenga (1 and 2 years,
344 respectively). The reason was ascribed to the fact that, after the beginning of cultivation, the farmer
345 has always practiced mulching in between and, with controlled crawling fires to reduce weeds. It
346 is possible that mulching of crop residues combined with the scarce fire application has increased
347 the organic matter content and reduced erosion so to maintain soil fertility that has supported the
348 crop yields for all the cultivated years. This practice resulted ineffective for the AvP, whose
349 generalized low values were attributed to the selective adsorption of phosphates on the abundant
350 Fe- or Al-oxyhydroxides (Parfitt, 1989; González-Rodríguez and Fernández-Marcos, 2018; Rafael
351 et al., 2020) and to the lack of fertilizer application. However, high contents of extractable Al and
352 Fe have also a positive aspect, being the Al- and Fe-oxyhydroxides particularly involved in the
353 stabilization of organic matter via the formation of complexes (Six et al., 2000; Verde et al., 2005;
354 Zhao et al., 2017; Totsche et al., 2018) and in the formation of stable soil structure. The higher
355 concentrations of organic matter in the soils of Macate, compared to Sussundenga and Vanduzi,
356 are likely related to their higher concentrations of dithionite-extractable iron and aluminum.

357 The soils of Macate also showed the highest TN content. Trees in the Macate FF were vigorous and
358 many belonged to the *Fabaceae* family, plants that are well-known for establishing symbiotic
359 associations with N₂-fixing bacteria (e.g., De Boer and Kowalchuk, 2001; Franche et al., 2009). Also,
360 the FF soils of Vanduzi and Sussundenga hosted leguminous plants, but they were less vigorous and
361 subjected to frequent stresses due to fire applications for hunting purposes or disturbances to obtain
362 firewood, timber, and charcoal. Being the FF of Macate older and less stressed, as demonstrated by

363 the presence of a rather well-expressed litter (Oi and Oe&Oa horizons), than those of Vanduzi and
364 Sussundenga, the relatively higher soil N content was ascribed to the presence of leguminous trees
365 that enriched the soil because of 50 years of undisturbed plant-microbial association. For the CF, even
366 though cultivated for 16 years with maize, the soil contained a relatively high TN content. Although
367 a possible contribution of N₂ fixation due to annual leguminous weeds like *Mucuna pruriens* cannot
368 be excluded, the relatively high TN content was ascribed to the continuous mulching (Fang et al.,
369 2011; Dong et al., 2018) and the scarce fire application. In fact, N is an extremely volatile element
370 that can be easily lost during the vegetation burning (Da Silva Neto et al., 2019).

371 In well-drained soils with acidic pH, low cation exchange capacity, and quartz dominated mineralogy
372 like those here studied, fertility is extremely low (Eshett et al., 1989), and proper soil management
373 including soil mulching and absence of fire is crucial to guarantee a minimum level of organic matter
374 and nutrient stocks, able to support crop yields (Bahr et al., 2014; Temudo et al., 2017).

375

376 4.1.3 Soluble anions and cations in soils

377 Well-drained and acid soils like those here studied are easily leached out of the most soluble ions
378 (Juo and Manu, 1996), so it appeared reasonable that the water-extracted anions and cations assumed
379 concentrations at the level of $\mu\text{eq kg}^{-1}$. Although in low concentrations, significant differences were
380 observed for chloride, fluoride, nitrate, and ammonium. Chloride was detected in higher contents in
381 the soil of Macate FF than in those of Vanduzi and Sussundenga. Among the several sources of
382 chloride listed by Geilfus (2019), since for Mozambique there is no report accounting for halite rock
383 outcrops except for a mixture of limestone and halite (Jofane formation) lying at 150-250 km from
384 the study sites (Schlüter, 2008), notwithstanding the long distance between study sites and the
385 Mozambique sea channel, the contribution of the airborne sea salts cannot be excluded. Especially at
386 Vanduzi, the angular coefficient of the Cl/Na relationship of the soil extract was similar to that of the
387 seawater (Keene et al., 1986) (Fig. S1 of Supplementary materials). Differences in fluoride content
388 can be due to natural and anthropogenic sources (e.g., Ali et al., 2016; Mikkonen et al. 2018; Wang

389 et al., 2019). Parent material is considered the most common source of fluoride, which is present in
390 minerals like fluorites, apatites, and micas, which form granite and igneous rocks, but also in topaz,
391 which is commonly associated with silicic igneous rocks, and cryolite, which comprises pegmatitic
392 rocks (e.g., Battaleb-Looie et al., 2012; Mikkonen et al., 2018). In the soils, fluoride was detected at
393 the highest concentrations under both CF and FF of Vanduzi and Macate, namely in soils with
394 different parent rocks; because of this, the geogenic source of fluoride was considered irrelevant in
395 determining the differences. Consequently, as largely reported in the literature (e.g., Feng et al., 2002;
396 Choubisa and Choubisa, 2016; Wang et al., 2019), the relatively higher fluoride content in the soils
397 of Vanduzi and Macate was ascribed to windblown materials coming from the surrounding mining
398 activities devoted to the extraction and transformation of copper, nickel, gold, silver, iron, and
399 bauxite, which are absent in a radius of at least 30 km from Sussundenga (Lehto and Gonçalves, 2008;
400 MIREME, 2021). In the Macate soils, where there was the highest content of TN, also nitrate and
401 ammonium were relatively abundant, possibly because of mulching. In fact, Yaşar Korkanç (2021)
402 demonstrated that soil mulching with organic materials reduced nitrate and ammonium losses through
403 water runoff when compared to uncovered plots and, for environments where rainstorms and cyclones
404 are frequent, mulching could make the difference. In the Sussundenga soils, where fertilizer has never
405 been used, only ammonium showed the highest concentrations possibly because of the ongoing
406 mineralization of tree roots after the recent (one year) slash and burn (Juo and Manu, 1996; Vitousek,
407 1981; Béliveau et al., 2015). In addition, ammonium is the direct product of biomass burning
408 (Knicker, 2007), and the recent slash and burn could have contributed to the ammonium release.
409 Therefore, concentrations of anions and cations were mainly ascribed to the arrival of windblown
410 materials for chloride and fluoride, and to the soil management for nitrate and ammonium, with no
411 relation to the different lengths of the forest-fallow period.

412

413 4.2 *Comparison of soils under forest-fallow and cropping within the location*

414 4.2.1 *Soil morphological and physical properties*

415 In each location, CF and FF soils were similar, and only at Macate, the FF soil contained a
416 minimum amount of 2:1 clay minerals, which were absent in the CF soil. Even though forest
417 vegetation is efficient in controlling soil erosion and runoff, especially when the understorey is thick
418 and made of shrubs and herbaceous vegetation (Doerr and Cerdà, 2005; Thomaz, 2013), it is also true
419 that in slash and burn systems, in addition to temperature fluctuations and direct rainfall, agronomic
420 practices like tillage can increase mineral weathering (Lemenih et al., 2005). Therefore, it seems
421 unrealistic that these differences in 2:1 clay minerals between FF and CF soils at Macate have been
422 produced in only 16 years of cultivation. Because of this, we considered this as a condition ascribable
423 to a site-specific difference of the parent material, which contained clay minerals or was more prone
424 to produce clay minerals through weathering.

425

426 4.2.2 *Soil chemical properties*

427 Significant differences in chemical properties between CF and FF were observed only for AvP in
428 the soils of Vanduzi. The slight P deficiency in the CF was probably exacerbated due to plant
429 absorption and erosion processes (Da Silva Neto et al., 2019), even though the CF was established
430 only one year before the soil sampling. In Oxisols, the richness of Fe-oxyhydroxides promotes a
431 strong immobilization of phosphates (Lü et al., 2017; Markovic et al., 2019), and the plant absorption
432 in absence of fertilizer application can rapidly decline the amount of AvP (Shen et al., 2011).
433 Considering the small chemical differences between CF and FF soils within locations, it appeared
434 that the effect of reforestation on the recovery of soil fertility was irrelevant.

435

436 4.2.3 *Soluble anions and cations in soils*

437 Only at Vanduzi the soil under CF appeared enriched with chloride and nitrate compared with that
438 under FF. In general, possible vectors of chloride and nitrate are irrigation water, fertilizers, and/or

439 other human activities (e.g., Geilfus 2019; Martinez Uribe et al., 2020; Oberhelman and Peterson,
440 2020). In this case, manual irrigation is practiced taking advantage of the water of a nearby stream
441 that, in the previous five km, crosses an area devoted to industrial poultry production, an urban area
442 at four km NW from the city of Chimoio, and a nearby industrial area. Concerning fertilization,
443 Geilfus (2019) mentioned as possible sources of chlorides animal wastes like cattle manure, pig
444 slurries, and chicken or pigeon manures. Both manure and sewage are possible sources of nitrates
445 (Jin et al., 2018; Wakida and Lerner, 2005; Torres-Martínez et al., 2021). Because of this, in the
446 absence of information about the water quality of the stream over the years, it is reasonable that the
447 relatively high content of chloride and nitrate in the CF soil of Vanduzi is mostly derived from the
448 irrigation water collected from the near stream. However, the CF is located in an educational
449 institution where, in addition to agriculture, extensive free-range cattle livestock is practiced during
450 the offseason, and cattle manure could be an additional source of nitrate and chloride.
451 Therefore, also for cations and anions distribution, soil management rather than the length of the
452 forest-fallow period appeared to exert some effect.

453

454 **5 Conclusions**

455 The CF and FF soils of the studied slash and burn system showed low physicochemical fertility,
456 with small differences depending mainly on the parent rock and soil management. Even if detailed
457 information about agroecological settings was impossible to obtain for this region, it was possible to
458 assess that FF soils were not in better conditions than those under CF and that the recovery of soil
459 fertility appeared inconsistent at least for a maximum of 50 years old forest-fallow, which represented
460 the longest forest-fallow period considered in this paper. Based on soil organic matter, TN, nitrate,
461 and ammonium contents, it seems that soil management such as mulching, irrigation, and reduced
462 fire application foster the recovery of the soils. It is therefore beneficial to adopt these agronomic
463 practices aimed at maintaining (or increasing) crop yields, but also at reducing the pressure on forest

464 land due to the frequent cuts and fires, with consequent improvement for macro- and micro- fauna
465 and flora biodiversity, at least in the first phase of the forest-fallow period.

466

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480

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